The background of the slide is a composite image. On the right side, a large, detailed view of a giant exoplanet is shown, featuring a mix of orange, red, and brownish-yellow colors, suggesting a hot, gaseous atmosphere. On the left side, a smaller, glowing protoplanet is depicted, surrounded by a cloud of gas and dust, representing an earlier stage of planetary formation. The background is a dark space filled with distant stars.

**Structure and evolution of (giant) exoplanets:
some news from the theoretical front**

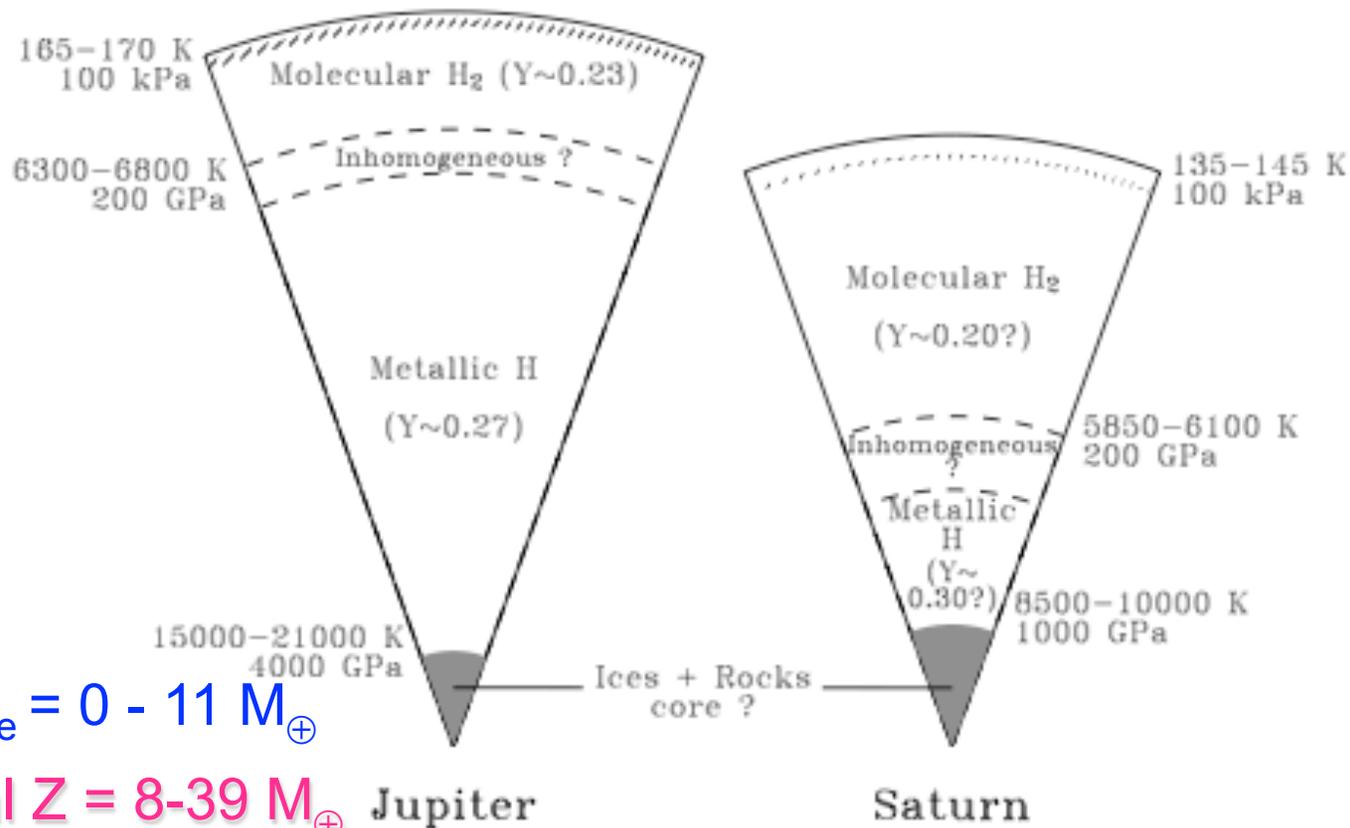
I. Baraffe
University of Exeter

I) Structure of Jupiter and Saturn

II) Exoplanets: Interior structure and evolutionary models:

- *Heavy element enrichment*
- *Inflated exoplanets*

Internal structure of Jupiter and Saturn *(the standard 3-layer picture)*



- $M_{\text{core}} = 0 - 11 M_{\oplus}$

- Total Z = 8-39 M_{\oplus}

(Saumon & Guillot 2004)

Recent picture

- **Small core hypothesis** for Jupiter **challenged** by recent calculations based on first-principles EOS for H-He mixtures (*Militzer et al. 2008*)

- ↳ based on a 2-layer model: core of rock/H₂O and isentropic mantle of H/He

- ⇒ Find a core for Jupiter of $M_{\text{core}} = 14\text{-}18 M_{\oplus}$

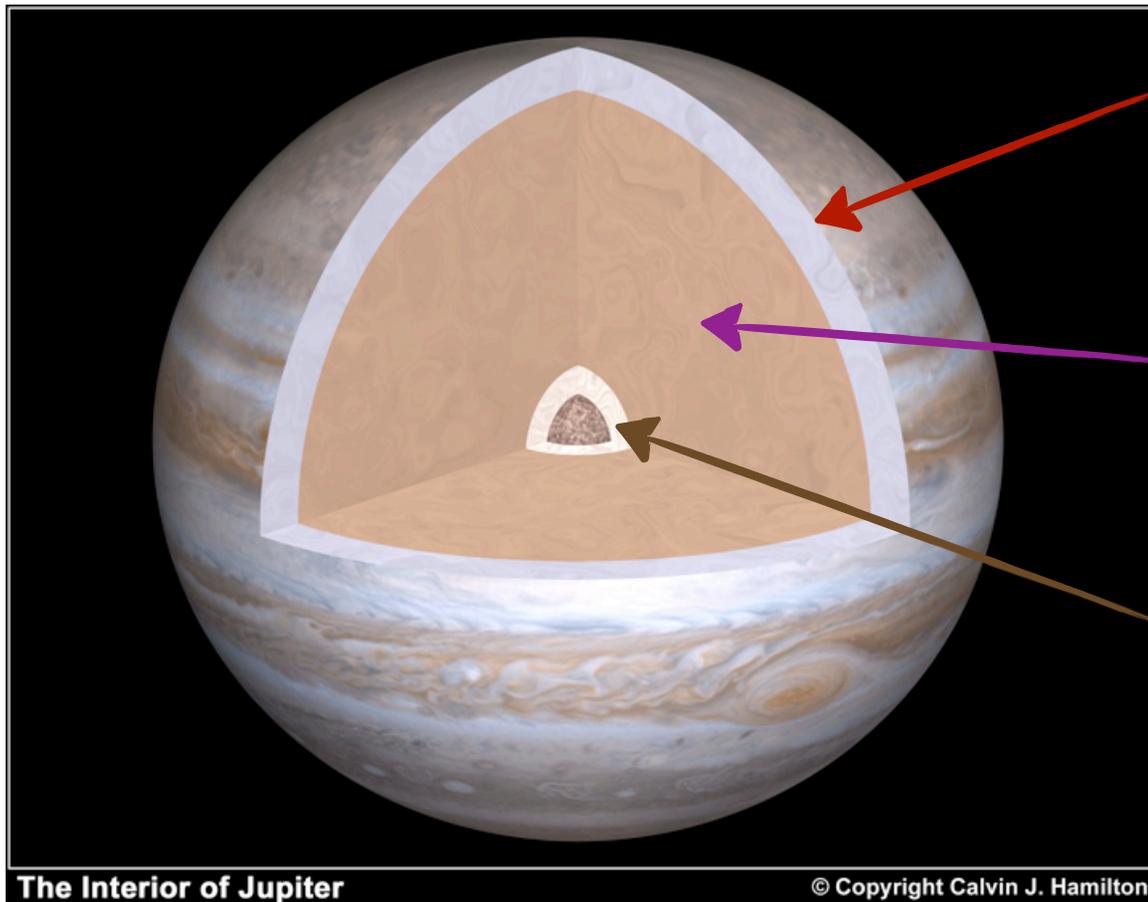
- **Strong disagreement** with another study also based on ab initio EOS calculations for H, He and H₂O (*Nettelmann et al. 2008*)

- ↳ based on a 3-layer model: core of rocks/ice + inner isentropic envelop (Metallic H, He, Z_{met}) + outer isentropic envelope (molecular H₂, He, Z_{mol})

- ⇒ Find a core for Jupiter of $M_{\text{core}} = 0\text{-}7 M_{\oplus}$, $M_Z = 30\text{-}50 M_{\oplus}$

II) Exoplanet modelling

(Baraffe et al. '06, '08, '10; Leconte et al. 2009, 10, 11)



The Interior of Jupiter

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Irradiated
atmospheres

*Barman et al.
2001;2005*

H/He envelope

$Z=0$ or $Z=0.02$

*Saumon &
Chabrier EOS*

Rocky/icy core

«Ices» (H_2O , CH_4 ,
 NH_3)

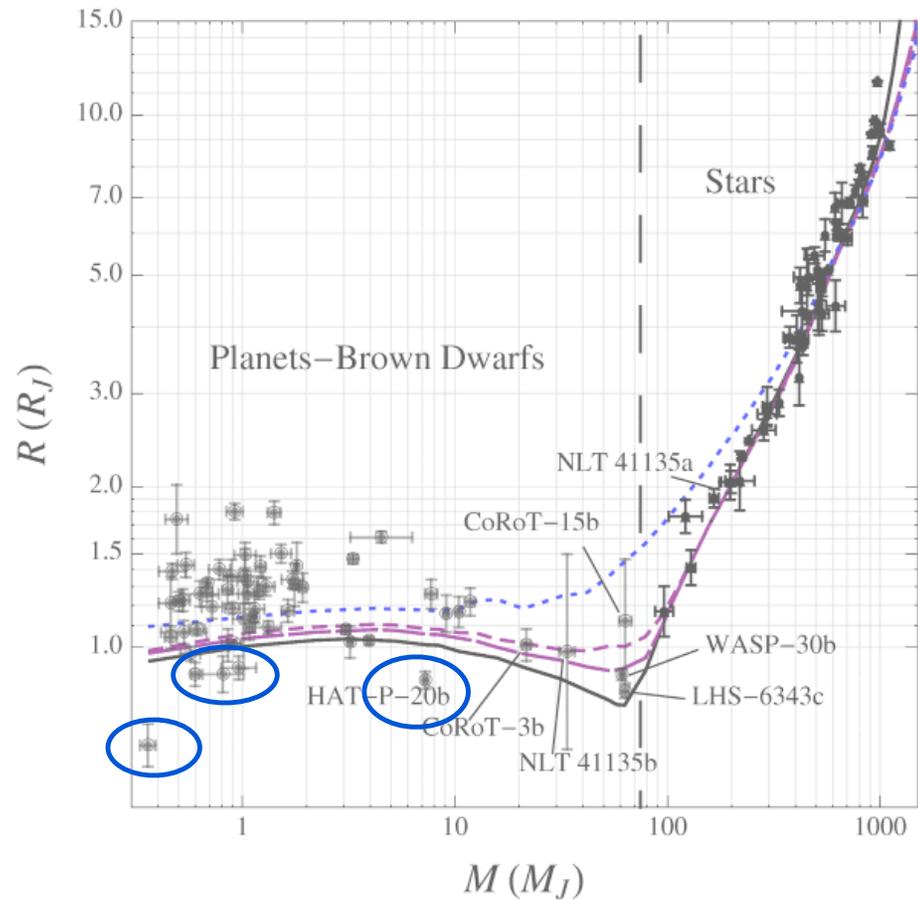
silicates ($MgSiO_4$,
 $MgSiO_3$,...)

Iron (Fe)

II-a) Heavy element enrichment

Substantial fraction of exoplanets are enriched in heavy material :

(volatiles H_2O , CH_4 , NH_3 ; silicates Mg_2SiO_4 , MgSiO_3 ; Fe)



Models of Baraffe et al. for Z_{\odot}
(0.1 - 1 - 5 - 10 Gyr)

- **Equations of state of heavy elements**

$10 M_{\oplus} \rightarrow 10 M_{\text{Jup}}$: extreme pressure and temperature regime

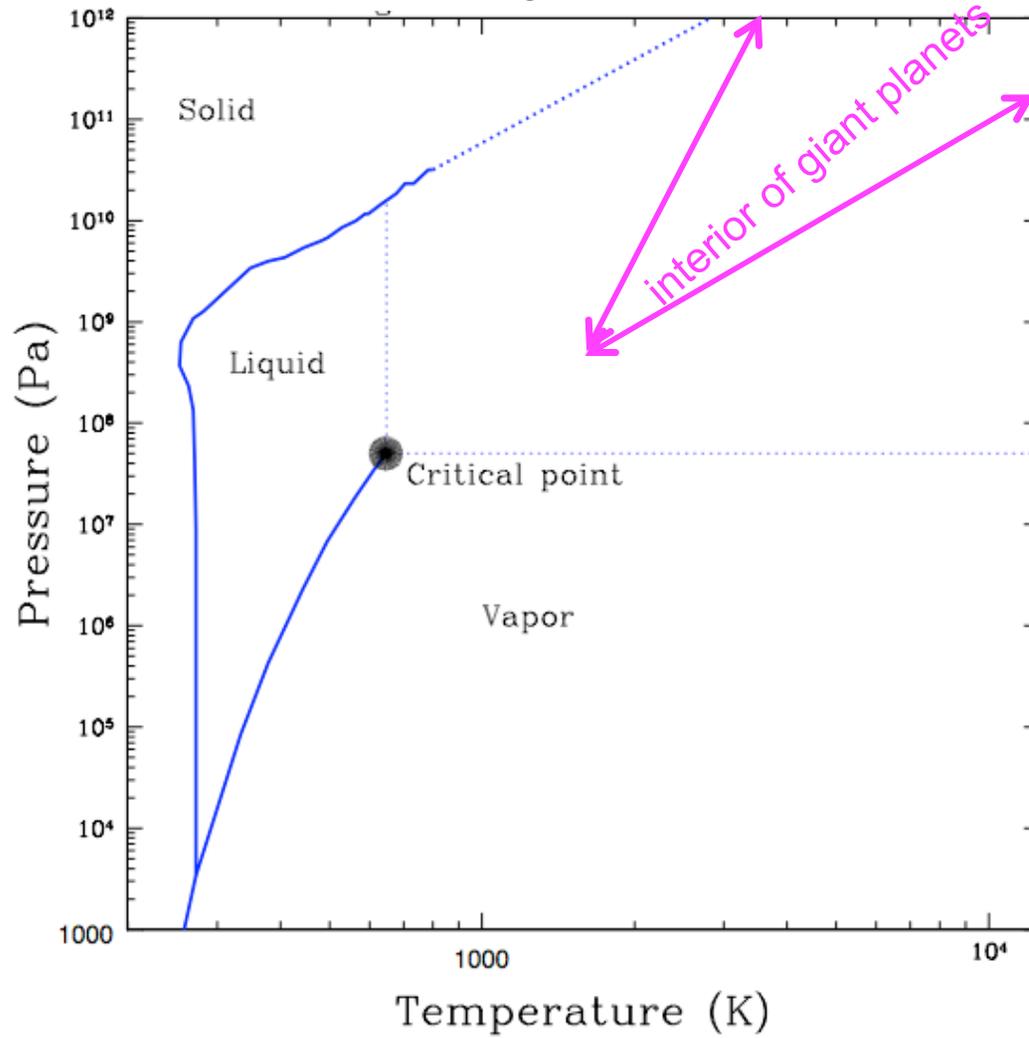
Mostly used EOS: ANEOS (Sandia) et SESAME (Los Alamos)

H_2O or « ice » (H_2O , CH_4 , NH_3)

Rock (MgSiO_4 , MgSiO_3 , etc..)

Fe

Phase diagram of water



▪ Distribution of heavy elements in planets

---> Current assumptions (Fortney et al. ; Burrows et al. etc...):

- All heavy elements located in the central core
- Metal-free or solar metallicity H/He envelope

Equivalent to a distribution of Z over the entire planet?

Z=50%

----> « all Z in the core » versus
« H/He/Z mixture in the entire planet »:
up to ~ 30 % effect on R at a given age

Baraffe, Chabrier, Barman 2008, 2010

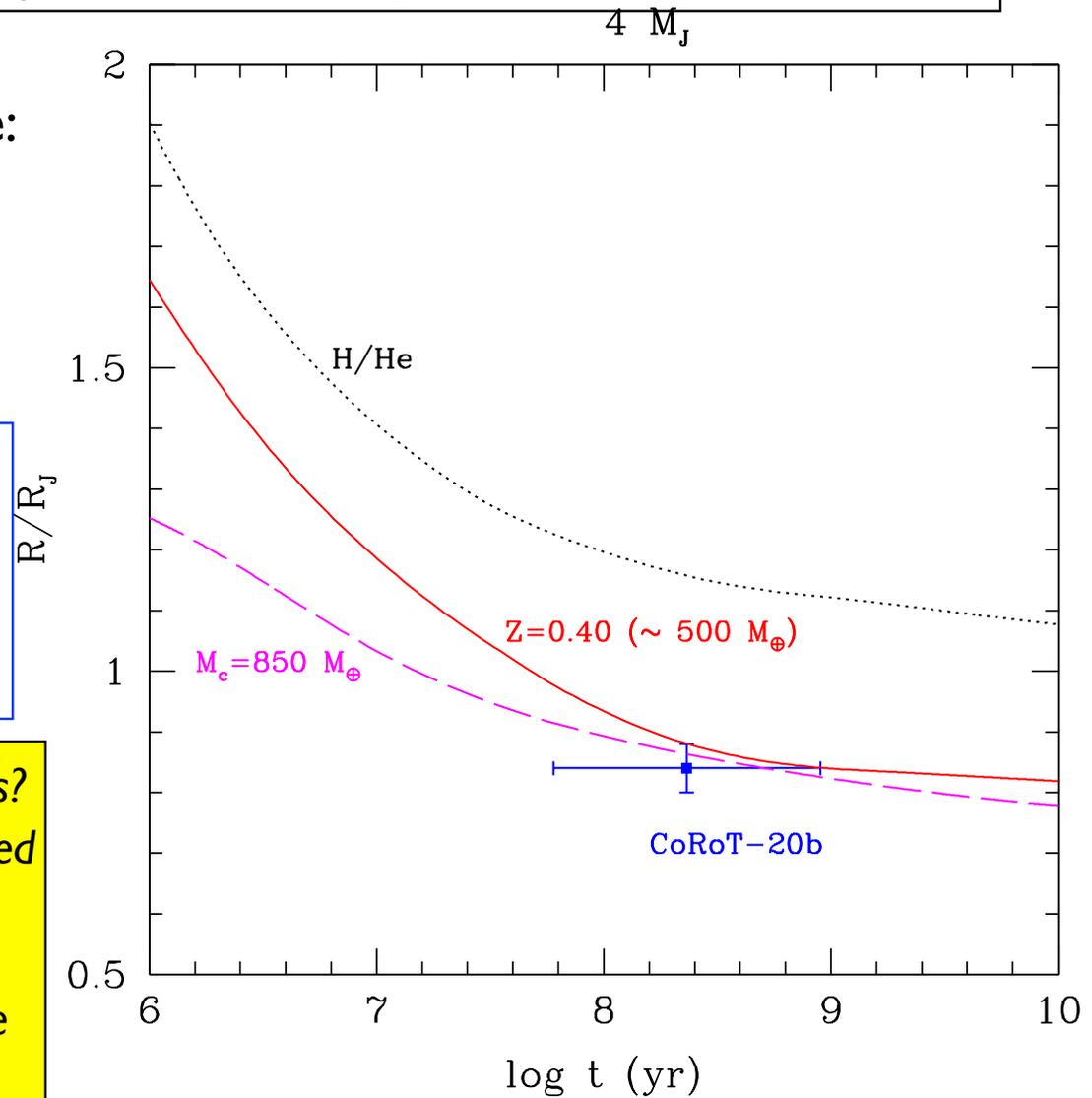
Bulk composition can provide a signature of the formation process (e.g. large heavy material content)

A very interesting case:
CoRoT20b
($4 M_{\text{Jup}}$ $0.8 R_{\text{Jup}}$)

Deleuil et al. 2012

Requires **too massive core** of heavy material to explain its radius (maximum amount of heavy material in the disk $\sim 800 M_{\text{earth}}$)

- Wrong estimate of radius?
- Heavy material distributed all over the planet?
- Pb with EOS used (could ab-initio EOS improve that?)

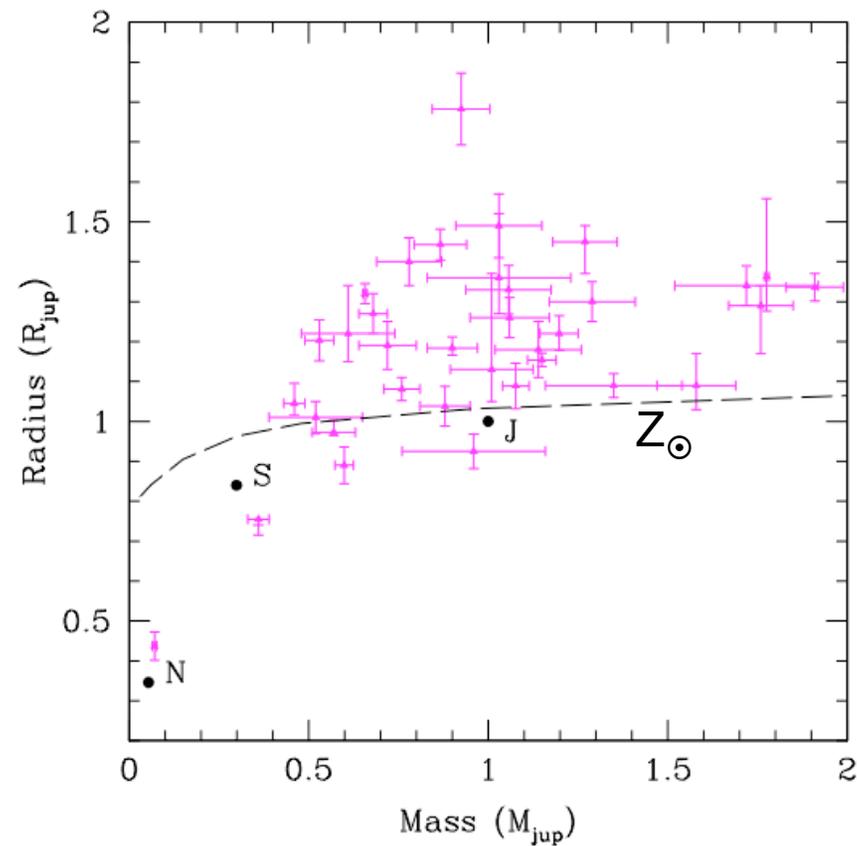


II-b) Inner structure models: inflated planets

Significant fraction of exoplanets with abnormally large radius

Missing physics in planetary interior models?

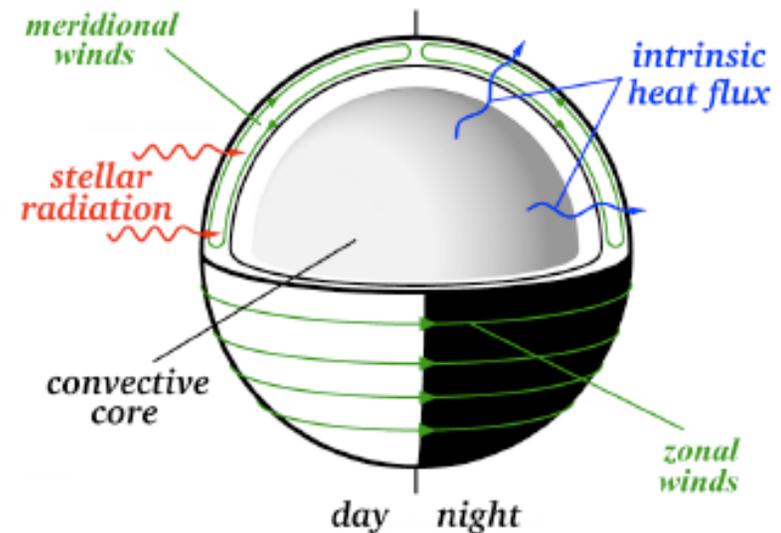
Several suggestions to explain this puzzle



a) **Atmospheric circulation:** Showman & Guillot 2002

----> irradiation from the parent star
creates strong **winds**

----> **downward** transport of **kinetic energy** down to the internal adiabat
→ *Heats the planet and slows down the contraction*



Atmospheric circulation models (GCM) (*Cho, Menou, Showman, etc....*)

➤ Adaptation of the **Exeter Met Office climate code** to exoplanets (can extend deeper than standard GCM codes, [see poster D. Acreman](#))

➤ **Study deep circulation pattern**

➤ **Effect of circulation on planet spectral signatures**
(*link with observations: HST, ECHO*)

b) Reduced heat transport: *Chabrier & Baraffe 2007*

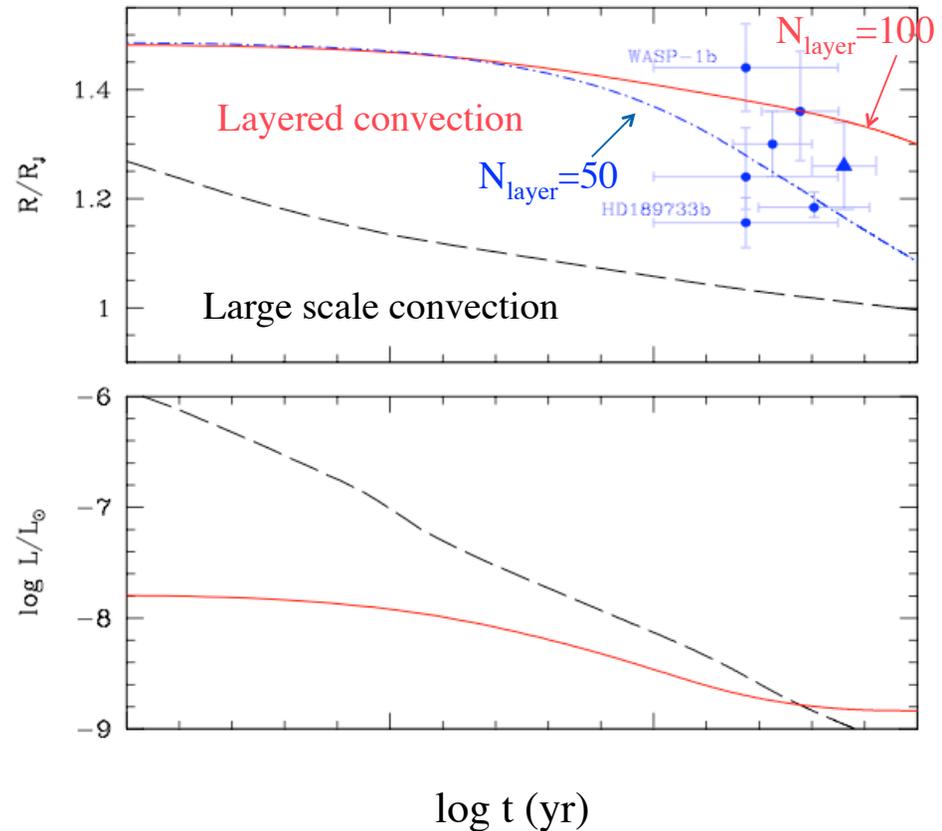
- Phenomenological approach
- **Idea:** reduced heat transport in planetary interior due to molecular weight gradient

⇒ « **layered convection** » : system of convective layers + thin diffusive layers
(double diffusive convection or semiconvection)

Main effect: **slow down** the evolution because of reduced heat transport in the interior

👉 planets with larger R

👉 Luminosity at young ages (< a few Gyr) **much lower** (testable with *Sphere*, *Gemini Planet Imager*)



1 M_J irradiated by a Sun at $a=0.05\text{AU}$

- **Double-diffusive convection in Jupiter and Saturn?**

(Leconte & Chabrier 2012, A&A, in press)

- ➡ Non conventional interior model for J and S
core + inhomogeneous, semiconvective envelope

- ➡ Reproduce the gravitational moments J_2 and J_4

- Jupiter: $M_{\text{core}} = 0 - 0.5 M_{\oplus}$

$M_{\text{env}}(\text{heavy}) = 41 - 64 M_{\oplus} \Rightarrow Z_{\text{tot}} = 13\% - 20\%$

- Saturne: $M_{\text{core}} = 12 - 21 M_{\oplus}$

$M_{\text{env}}(\text{heavy}) = 10 - 24 M_{\oplus} \Rightarrow Z_{\text{tot}} = 28\% - 44\%$

Inhomogeneous models for Jupiter and Saturn are significantly more enriched in heavy material (30%-60% more) than previously thought ➡ change our standard picture of “homogeneous” layers

Rosenblum, Garaud, Traxler, Stellmach 2011: 3D numerical simulations

→ Layers can form in low-Pr (< 1) double diffusive convection

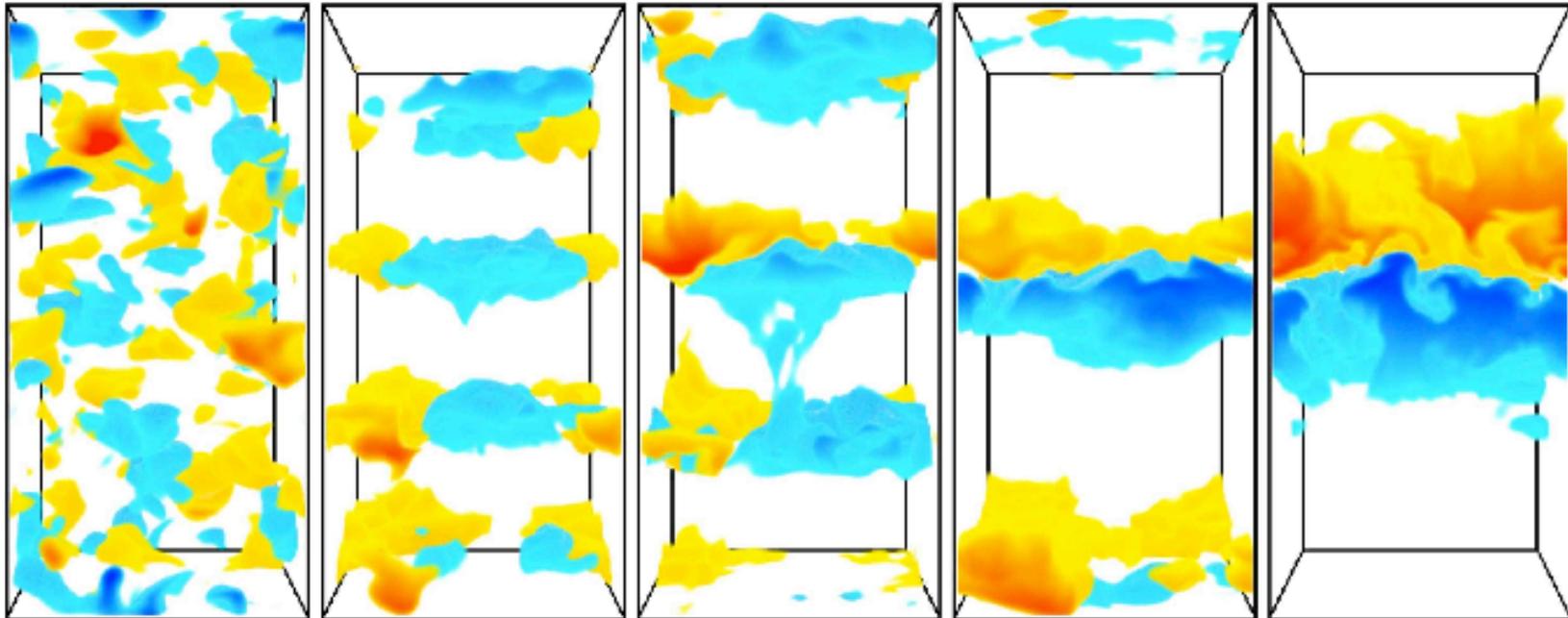


Fig. 5.— Volume-rendered visualization of the mean molecular weight perturbation, for

Question: do all layers merge or do the mergers stop and equilibrium layers form with height \ll size of system??

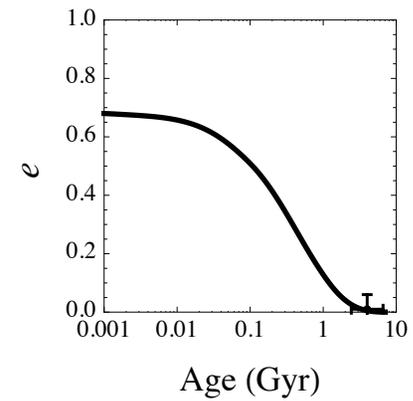
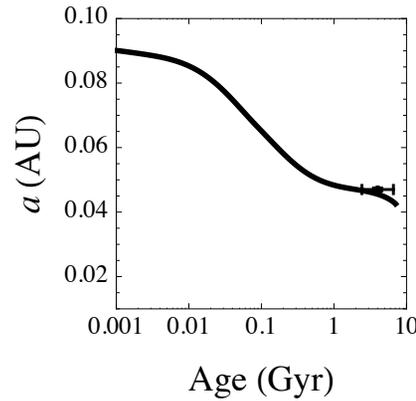
c) Tidal heating

(Bodenheimer et al. 01; Jackson et al. 08; Miller et al. 2009; Leconte et al. 2010)

HD 209458b

Leconte et al. 2010

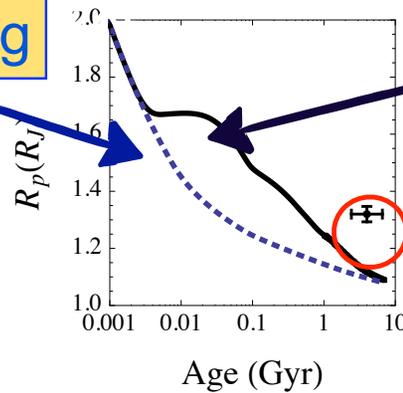
(tidal model non truncated in eccentricity)



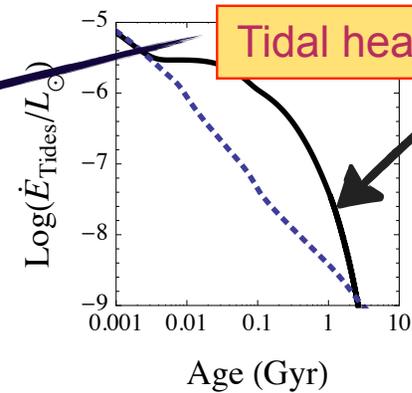
Semi-major axis

Eccentricity

Without heating



Tidal heating



Planetary radius

Tidal energy dissipation

➡ Difficult to explain with tidal effects alone properties of most inflated planets

The future: *(some future developments)*

- **Improved EOS** of H/He and heavy materials (water, silicates, etc) at high pressure and high temperature

Progress are coming!

- Ongoing and future high-pressure experiments (Livermore, Sandia in the US; LIL and Laser Megajoule in France)
 - First principle numerical methods (DFT, path integral, quantum molecular dynamics)
-
- Development of **numerical simulations** to confirm the existence and stability of **layered convection** in planetary interiors (*Rosenblum et al. 2011; Mirouh et al. 2012*)
 - ☞ Planets are not necessarily fully adiabatic and homogeneous
 - ☞ Important impact on our own giant planets!
-
- Development of **dynamical atmospheric models** (heating/cooling + circulation + radiative transfer)
 - ☞ Solution for abnormally large radii of close-in planets?
 - ☞ Effect on spectral signatures